

Deformation on plate boundaries; What it looks like, how we observe it and how we can correct coordinates for its effect: the example of the HTDP software and NAD83.

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Christchurch, New Zealand, 1-2 May 2016

NAD83(PACP00) aligned with Pacific Plate



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THE TRIGGERING MECHANISM: SLIPPAGE ALONG A FAULT



CRUSTAL BLOCKS AT REST



DEFORMATION DURING STRESS BUILD-UP



THE INSTANT OF RUPTURE

REBOUNDING TO A NEW EQUILIBRIUM





Time series for PYGR

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Secular velocity field for Western con US



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Secular velocity field for Western con US

- Western US comprised of independently rotating blocks bordered by faults
- Each segment of a fault has a locking coefficient
- Model solves for the Euler pole of each block plus the locking coefficients
- The velocity model is constrained by
 - 4643 GPS vectors
 - 166 slip vector/fault azimuth data
 - 170 fault slip rate measurements





FIG/IAG/UN-GGIM- AP/UN-ICG/NZIS Technical Seminar Reference Frame in Practice Christchurch, New Zealand, 1-2 May 2016

Profiles

Sponsors:

- Profiles 1-3 cross the part of the plate boundary dominated by the San Andreas Fault
- Profile 4 crosses the Cascadia Subduction zone

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Secular velocity across the plate boundary in North America





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Western CONUS grids

Grids

- *HTDP* uses interpolated grid files to calculate the secular velocities.
- These grids cover different regions with different cell sizes in order to obtain higher accuracy in regions of higher velocity gradients

Longitude range	Latitude range	Cell spacing	Region
West	North	(minutes)	
125° to 100°	31°-49°	15	Western CONUS
125° to 122°	40°-49°	3.75	Pacific NW
125° to 119°	36°-40°	3.75	Northern CA
121º to 114º	31°-36°	3.75	Southern CA
120.51° - 121.8°	35.8° – 36.79	0.6	San Andreas

How earthquakes are modeled



FIG/IAG/UN-GGIM- AP/UN-ICG/NZIS Technical Seminar **Reference Frame in Practice** Christchurch, New Zealand, 1-2 May 2016

Dislocations





- Each dislocation represents a rectangular patch where one side slips relative to the other
- The name, dislocation, is used because the slip displacement is uniform over the rectangle
- thereby producing a discontinuity along the edges of the patch.



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Post-seismic relaxation - 2009 Dusky Sound earthquake





Post seismic relexation

• HTDP uses an exponential decay function to model postseismic decay for the 2002 Denali Earthquake

$$D_{j}(\varphi, \lambda, t) = A_{j}(\varphi, \lambda) \left[1.0 - e^{\left(-\frac{(t-\tau)}{\nu}\right)} \right] \text{if } \tau < t$$
$$D_{j}(\varphi, \lambda, t) = 0 \text{ if } \tau > t$$

• Where v=5 years and $\tau=2000.841$



Post-seismic coefficients



167° 168° 169° 170° 171° 172° –43° **-4**4° –45° –46° –47°

Post-seismic deformation over 5 years





Velocity change



HTDP, v 3.0 included dislocation model for the 2002 Denali earthquake



Source: Elliott, J. L., Freymueller J. T., and Rabus B. (2007), Coseismic deformation of the 2002 Denali fault earthquake: Contributions from synthetic aperture radar range offsets, *J. Geophys. Res.*, 112, B06421, doi:10.1029/2006JB004428.



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Vertical velocities in New Zealand

Sigrun Herinsdottir GNS Science



InSAR derived velocity maps from 2003 – 2010 reveal a number of subsiding regions around New Zealand.

Subsidence of up to 10 mm/yr is observed across the Waihi Township. Most likely as a result of past mining activities



Ian Hamling GNS Science



InSAR derived velocity maps from 2003 – 2010 reveal a number of subsiding regions around New Zealand.

Widespread subsidence has been observed along the Taupo Volcanic Zone associated with both geothermal activities and cooling of magma at depth



Kawerau